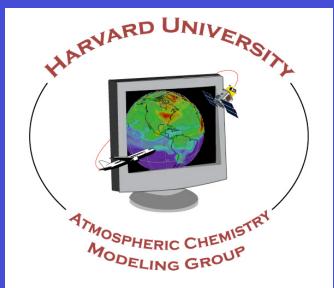
## CONSTRAINING AEROSOL SOURCES AND PROCESSES USING FIELD OBSERVATIONS AND MODELS

Daniel J. Jacob

with Tzung-May Fu<sup>1</sup>, Jun Wang<sup>2</sup>, Easan E. Drury<sup>3</sup>



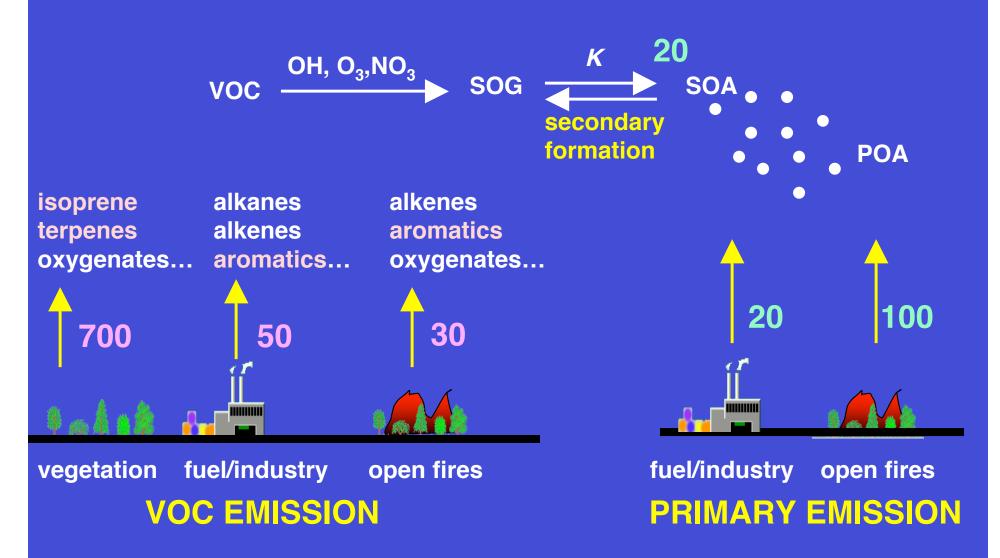
and funding from EPRI, NSF, NOAA, NASA

<sup>&</sup>lt;sup>1</sup> now asst. prof. at Honk Kong Polytechnic University

<sup>&</sup>lt;sup>2</sup> now asst. prof. at University of Nebraska

<sup>&</sup>lt;sup>3</sup> still trying to get out

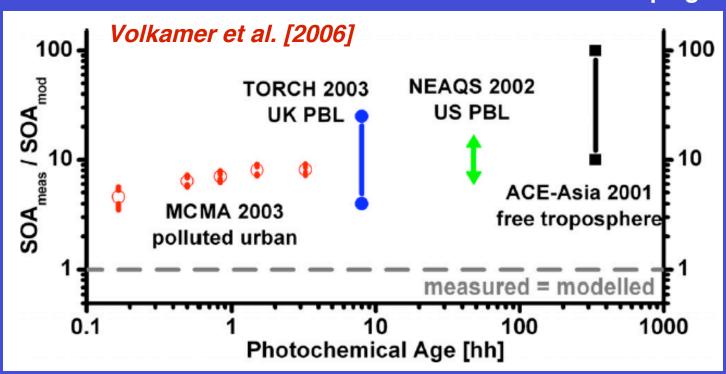
### CONVENTIONAL MODELING OF ORGANIC AEROSOL



Global sources in Tg C y<sup>-1</sup> (standard version of GEOS-Chem model)

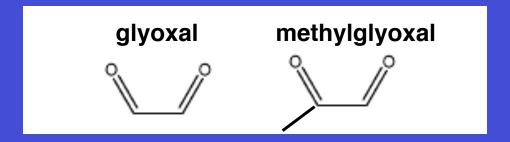
### ...BUT THESE MODELS UNDERESTIMATE OBSERVATIONS

simulated/observed ratios from recent measurement campaigns



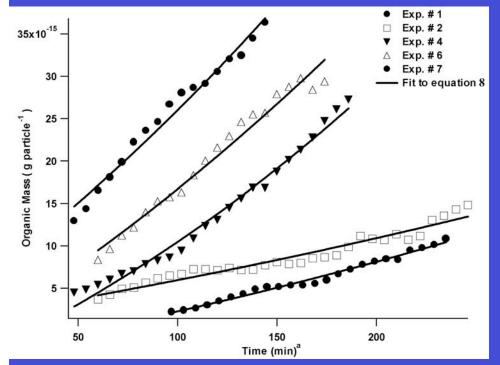
Discrepancy worsens as air masses age; suggests irreversible SOA source missing from the models

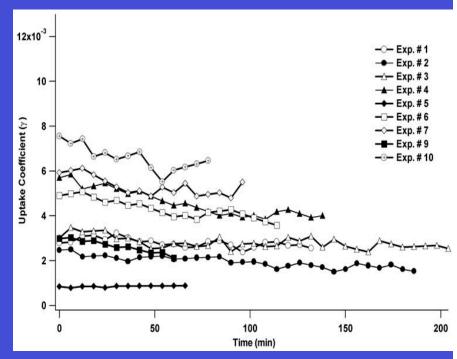
#### IRREVERSIBLE DICARBONYL UPTAKE BY AQUEOUS AEROSOL



Chamber AMS experiments of glyoxal uptake by Liggio et al. [JGR 2005] Inferred reactive uptake coefficient y

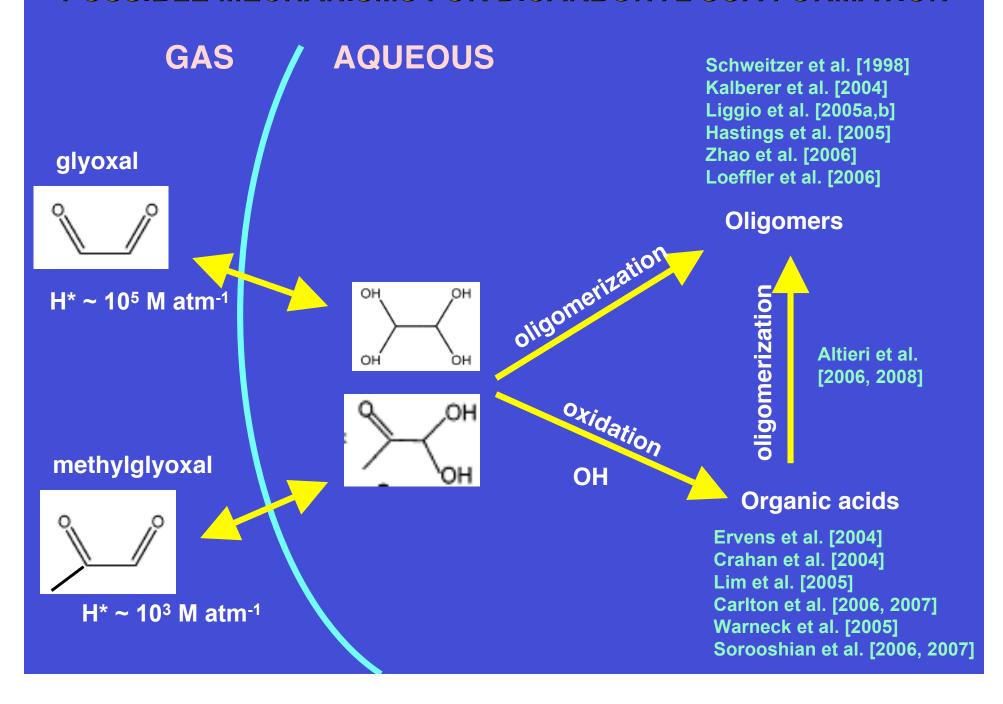
Organic aerosol mass growth with time



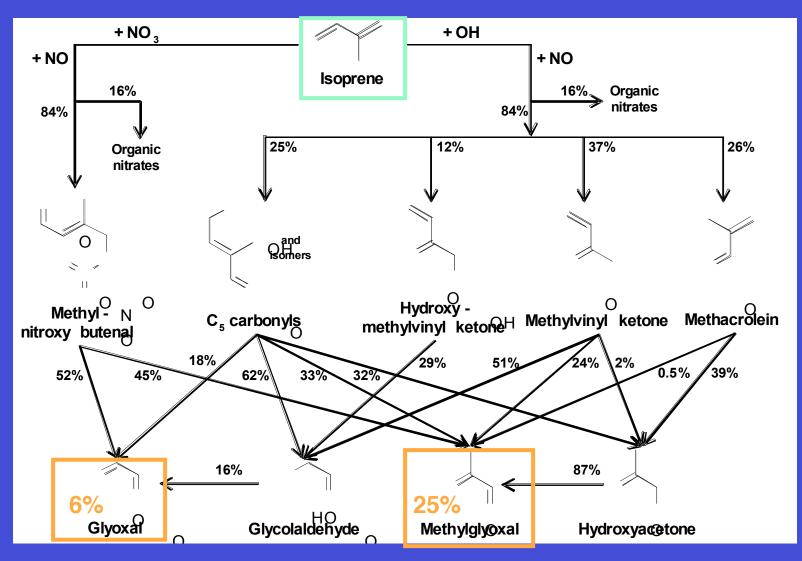


- median  $\gamma = 2.9 \times 10^{-3}$  observed for aqueous surfaces; evidence for oligomerization
- similar γ observed for methylglyoxal on acidic surfaces [Zhao et al. ES&T 2006]

### POSSIBLE MECHANISMS FOR DICARBONYL SOA FORMATION



### **GLYOXAL/METHYLGLYOXAL FORMATION FROM ISOPRENE**



GEOS-Chem mechanism based on MCM v3.1

### **GLOBAL GLYOXAL BUDGET IN GEOS-Chem**

Including reactive uptake by aq. aerosols + clouds with  $\gamma = 2.9 \times 10^{-3}$  [Liggio et al., 2005]

СНОСНО			
Production E	mission [Tg y <sup>-1</sup> ]	Molar yield [%]	45 [Tg y <sup>-1</sup> ]
Isoprene	410	6.2	21
Acetylene	6.3	64	8.9
Glyoxal †biomass burni	ng) 7.7	100	7.7
Ethylene	21	5.7	2.5
Monoterpenes	160	2.8	1.8
Benzene	4.8	25	0.9
Toluene	6.7	16	0.7
Xylenes	4.7	16	0.4
Glycolaldehyde *	5.6	9.9	0.5
Methylbutenol	9.6	5.4	0.3
Loss			45 [Tg y <sup>-1</sup> ]
2.9 h Photolysis			28
Oxidation by OH			6.5
SOA formation			6.4
Dry deposition			2.2
Wet deposition			1.9

Global SOA formation of 6.4 Tg yr<sup>-1</sup> (1.0 in clear sky + 5.4 in cloud); compare to 16 Tg yr<sup>-1</sup> from terpenes/isoprene by semivolatile mechanism

### GLOBAL METHYLGLYOXAL BUDGET IN GEOS-Chem

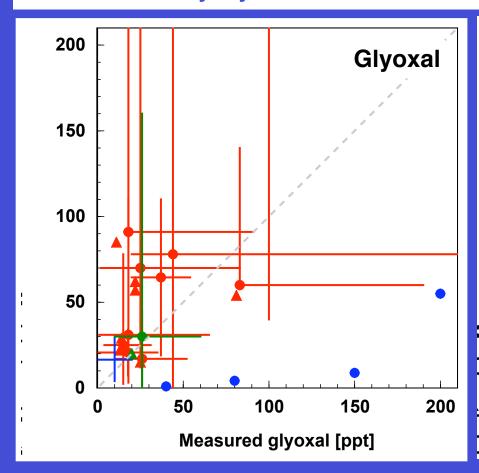
Including reactive uptake by aerosols and clouds with  $\gamma = 2.9 \times 10^{-3}$ 

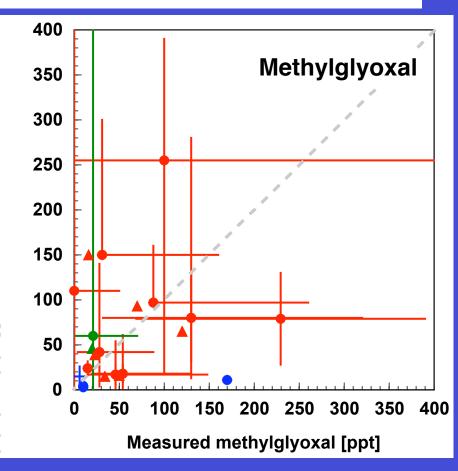
CH₃COCHO			
<b>Production</b>	Emission [Tg y <sup>-1</sup> ]	Molar yield [%]	140 [Tg y <sup>-1</sup> ]
Isoprene	410	25	110
Acetone	57	14	10
Methylglyoxal*(bioma	ass 5.0	100	5.0
>C <sub>2</sub> alkenes burnin	<sup>ig)</sup> 31	7.7	4.1
Hydroxyacetone*	4.9	<b>75</b>	3.6
Monoterpenes	160	4.2	3.5
Propane <sup>-</sup>	16	11	2.7
>C3 alkanes	26	3.2	1.0
Toluene	6.7	12	0.7
Xylenes	4.7	23	0.7
Methylbutenol	9.6	6.2	0.5
Loss			140 [Tg y <sup>-1</sup> ]
= 1.6 h Photolysis			100
SOA formation			16
Oxidation by OH			15
Wet deposition			1.8
Dry deposition			1.7

Global SOA formation of 16 Tg yr<sup>-1</sup> (2 in clear sky + 14 in cloud); compare to 16 Tg yr<sup>-1</sup> from terpenes/isoprene by semivolatile mechanism

### MODEL COMPARISON TO IN SITU OBSERVATIONS

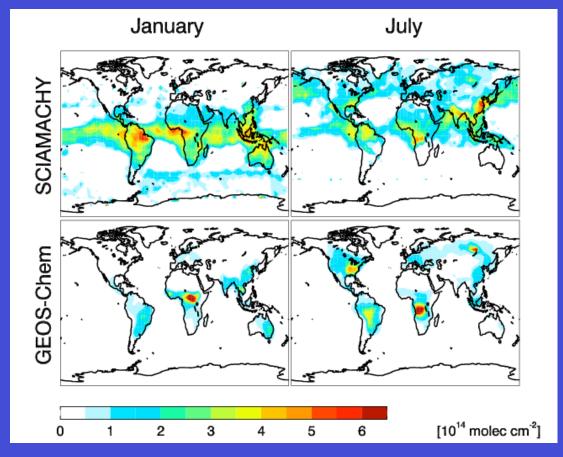
Continental boundary layer (all northern midlatitudes summer)
Continental free troposphere
Marine boundary layer





Indication of a missing marine source in the model

### SCIAMACHY SATELLITE OBSERVATION OF GLYOXAL

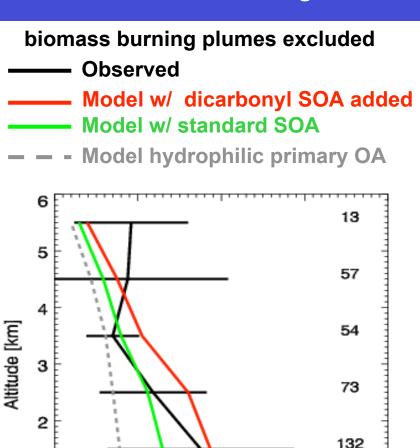


- General spatial pattern reproduced over land, SCIAMACHY is 50% higher than model
- SCIAMACHY sees high values over oceans correlated with chlorophyll: unidentified marine source?

100 pptv glyoxal in marine boundary layer would yield ~1 μg C m<sup>-3</sup> SOA; could contribute to observed OC aerosol concentrations in marine air Fu et al. [JGR, submited]

#### SIMULATION OF WSOC AEROSOL OVER EASTERN U.S.

Water-soluble OC (WSOC) aerosol observations by Rodney Weber (GIT) from NOAA aircraft during ICARTT campaign out of Portsmouth, NH (Jul-Aug 04)

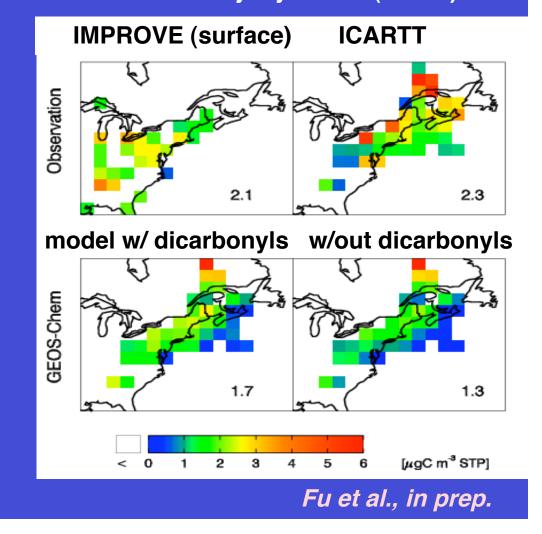


WSOC [µgC m<sup>-3</sup> STP]

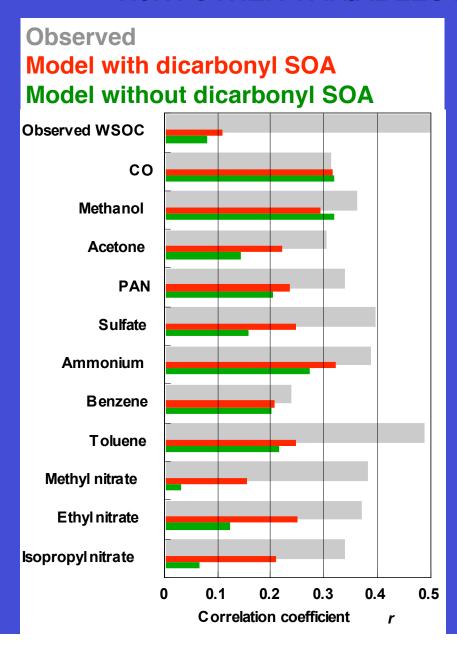
0

391

#### **Boundary layer data (<2 km)**



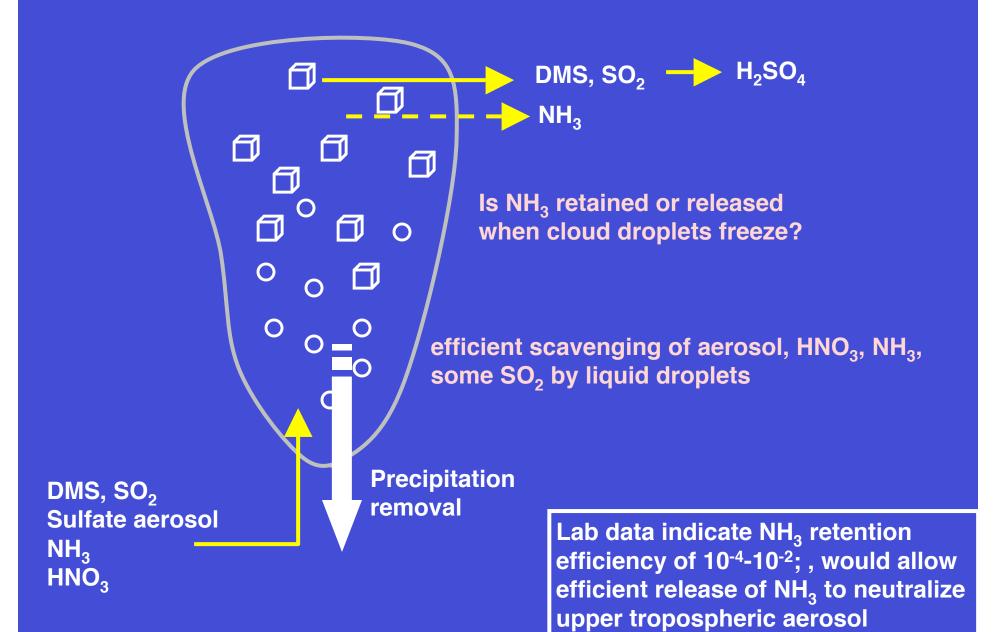
## CORRELATIONS OF FREE TROPOSPHERIC WSOC WITH OTHER VARIABLES MEASURED ON NOAA AIRCRAFT



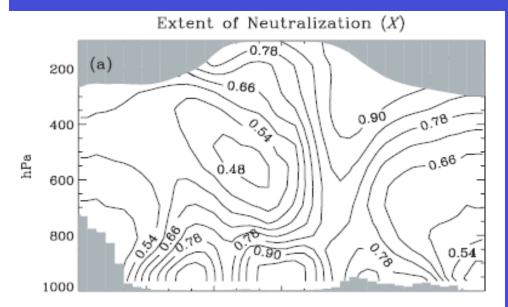
- WSOC is observed to correlate with
- toluene and methanol (anthro+bio?)
- sulfate (aqueous-phase production?)
- alkyl nitrates (photochemistry?)
- Model does not reproduce observed WSOC variability but does better with correlations, particularly when dicarbonyl SOA is included (sulfate, alkyl nitrates)

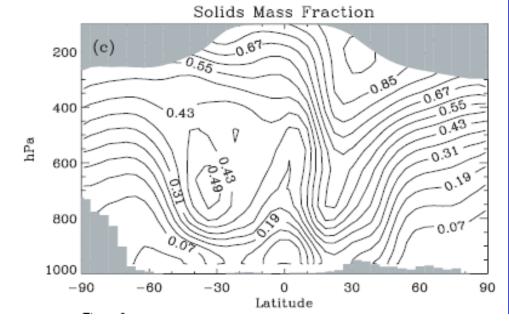
Fu et al., in prep.

# EXPLAINING PERSISTENT OBSERVATIONS OF NEUTRALIZED SULFATE IN UPPER TROPOSPHERE



# IMPLICATIONS FOR SULFATE NEUTRALIZED FRACTION (X) AND AEROSOL PHASE





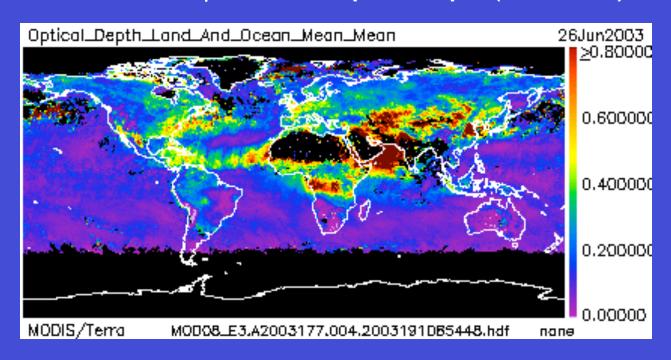
Annual zonal mean GEOS-Chem model results in an ammonium-sulfate simulation including hysteresis of phase transitions and NH<sub>3</sub> retention efficiency of 0.05 upon cloud freezing

Upper tropospheric sulfate is mostly neutralized and solid! Implications for atmospheric chemistry, cirrus formation...

Wang et al. [JGR, submitted]

# INTERPRETING SATELLITE AEROSOL DATA: HOW DO WE GO BEYOND PRETTY PICTURES?

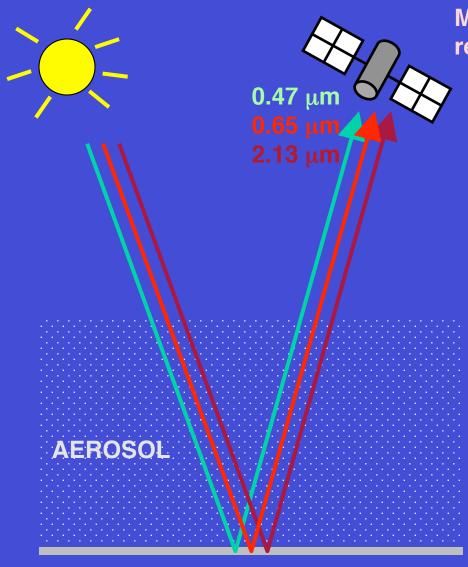
MODIS 0.47 μm aerosol optical depth (June 2003)



How can we use satellite data to better quantify aerosol sources and processes through comparison to models? Need

- 1. improved surface reflectance data over land
- 2. model simulation of top-of-atmosphere reflectance in satellite field of view

# IMPROVING MODIS SATELLITE RETRIEVALS OF AEROSOL OPTICAL DEPTHS OVER LAND



MODIS measures top-of-atmosphere (TOA) reflectance in several wavelength channels

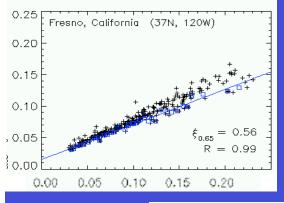
- Interpretation of TOA reflectance in terms of AOD requires assumptions on surface reflectance, aerosol optical properties
- Use TOA reflectance at 2.13 μm (transparent atmosphere) to derive surface reflectance
- MODIS operational algorithm relies on general assumptions for 0.47/2.13 and 0.65/2.13 surface reflectance ratios; we improve by deriving those locally using lower envelope in scatterplots of 0.65 vs. 2.13 MODIS TOA reflectance data
- MODIS operational algorithm relies on general categories for aerosol optical properties; improve by using local GEOS-Chem model data

**SURFACE** 

Drury et al. [JGR, subnmitted]

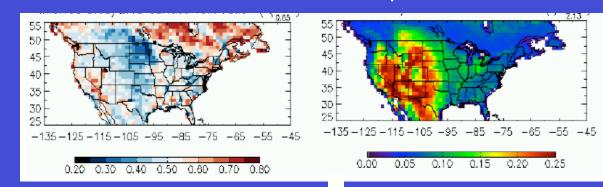
## **GEOS-Chem SIMULATION OF MODIS TOP-OF-ATMOSPHERE REFLECTANCE (JUL-AUG 2004)**

0.65 vs. 2.13 μm **TOA reflectance** 



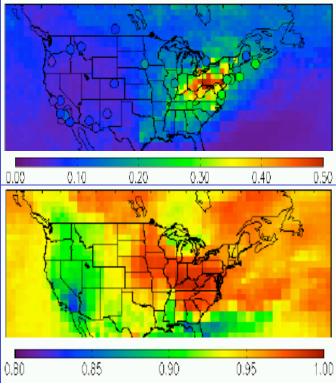
**0.65/2.13 surface reflectance ratio 2.13 μm TOA reflectance** 

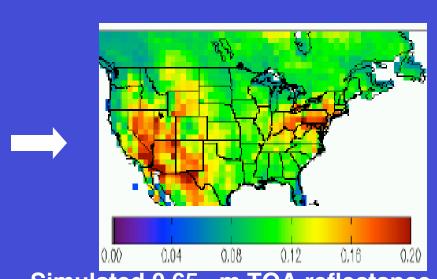




**GEOS-**Chem **0.65** μm **AOD** (AERONET In circles)

**GEOS-**Chem **0.65** μm singlescattering albedo





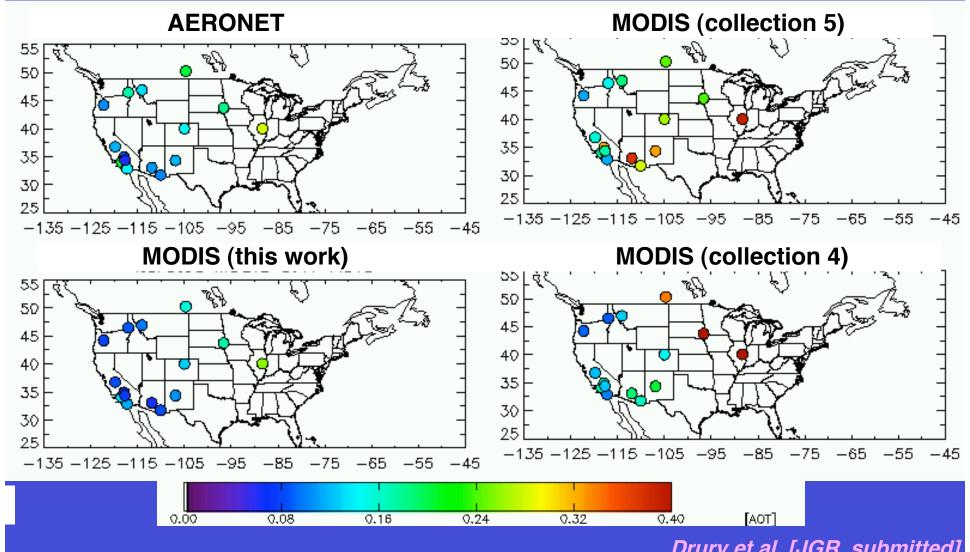
Simulated 0.65 µm TOA reflectance

Drury et al. [JGR, submitted]

### IMPROVED AOD RETRIEVAL OVER CENTRAL/WESTERN U.S.

by fitting model TOA reflectances to MODIS observations

MODIS vs. AERONET 0.47 μm AODs (Jul-Aug 2004)



Drury et al. [JGR, submitted]

#### NASA/ARCTAS 2008 AIRCRAFT CAMPAIGN TO THE ARCTIC

Two deployments: April (Fairbanks) and June-July (Cold Lake, Alberta)

Four research themes: (1) transport of mid-latitudes pollution to Arctic, (2) boreal forest fires, (3) aerosol radiative forcing, (4) chemical processes



### DC-8: in situ chemistry and aerosols

Ceiling 37 kft, range 4000 nmi, endurance 9 h Payload: O<sub>3</sub>, H<sub>2</sub>O, CO, CO<sub>2</sub>, CH<sub>4</sub>, NO<sub>x</sub> and HO<sub>x</sub> chemistry, BrO, mercury, NMVOCs, halocarbons, SO<sub>2</sub>. HCN/CH<sub>3</sub>CN, actinic fluxes, aerosol composition, aerosol mass and number concentrations, aerosol physical and optical properties, remote ozone and aerosol



#### P-3: radiation and in situ aerosols

Ceiling 30 kft, range 3800 nmi, endurance 8 h Payload: optical depth, radiative flux, radiance spectra, aerosol composition, black carbon

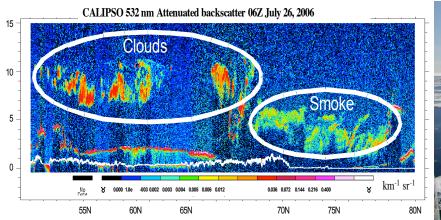


# B-200: aerosol remote sensing and CALIPSO validation

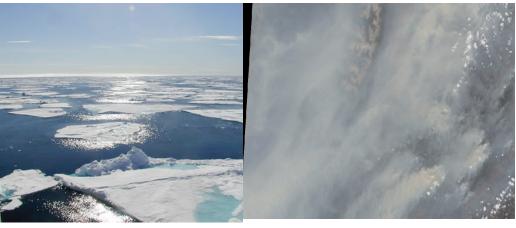
Ceiling 32 kft, range 800 nmi, endurance 3.5 h Payload: High Spectral Resolution Lidar (HSRL) Research Scanning Polarimeter (RSP)

### **ARCTAS Science Theme 3: Aerosol radiative forcing**

#### CALIPSO clouds and smoke



**Arctic haze** MISR true-color fire plume



C. Trepte, LaRC

R. Kahn, JPL

- What is the regional radiative forcing from Arctic haze, fire plumes?
- How does this forcing evolve during plume aging?
- What are the major sources of soot to the Arctic?
- How does soot deposition affect ice albedo?

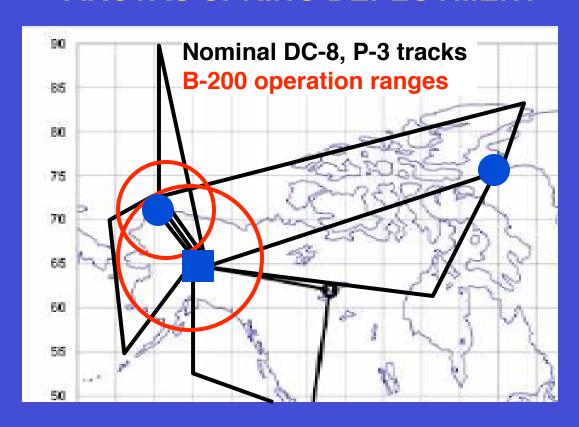
#### Satellite capabilities:

- UV/Vis/IR reflectances (Cloudsat, MODIS, MISR, OMI)
- multi-angle sensing (MISR)
- · lidar (CALIPSO)

#### Aircraft added value:

- detailed in situ aerosol characterization
- remote sensing of radiances, fluxes
- BRDFs

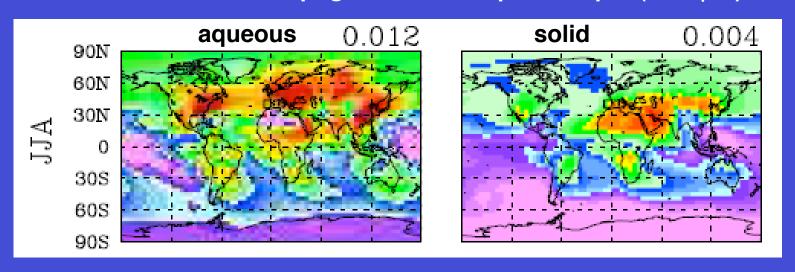
### **ARCTAS SPRING DEPLOYMENT**



- Deployment period: April 1-21
- About 70 flight hours for each aircraft
- Primary base: Fairbanks. Secondary bases: Barrow (B-200), Thule (DC-8, P-3)
- Several flights to involve collaboration with ISDAC

# EFFECT OF PHASE TRANSITIONS ON DIRECT SULFATE RADIATIVE FORCING

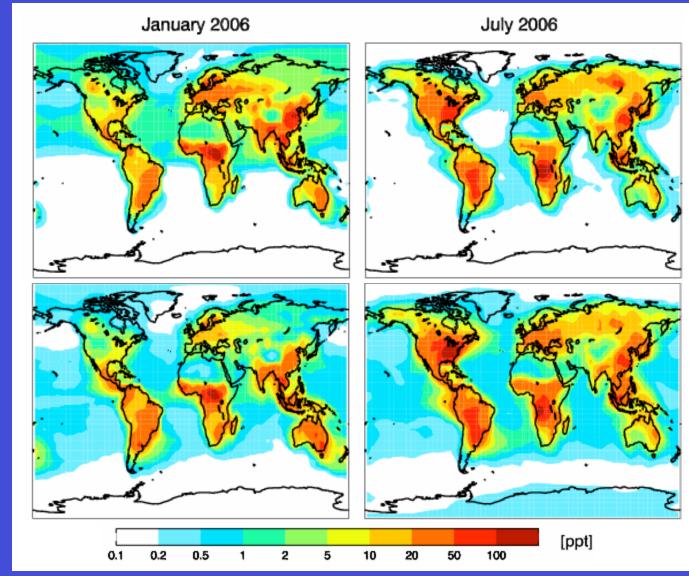
GEOS-Chem anthropogenic sulfate optical depth (0.55 μm)



Direct anthropogenic sulfate radiative forcing (W m<sup>-2</sup>)

Base simulation with hysteresis	-0.25
CRH=DRH (upper branch of hysteresis loop)	-0.24
DRH=CRH (lower branch of hysteresis look)	-0.26
CRH=0 (no crystallization)	-0.27

## SIMULATED SURFACE CONCENTRATIONS



- · Highest concentrations in regions of biomass burning regions, active vegetation
- ~1 ppt background from acetylene (glyoxal), acetone (methylglyoxal)

Glyoxal

Methyl-

glyoxal

